# Development of Mercury Modeling Schemes Within CMAQ Framework: Science and Model Implementation Issues

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# 1. INTRODUCTION

Mercury (Hg) is a persistent, bioaccumulative pollutant regulated by the United States Environmental Protection Agency (EPA). The concern of mercury pollution arises from the health effects caused by mercury vapor inhalation through the contaminated air, and methylated mercury ingestion through the consumption of fresh water and marine fish. According to estimates of the global mercury budget, about 6.000 tons and 10,800 tons of mercury are currently present in the troposphere and in the water bodies on earth. Mercury is released into the atmosphere from a variety of natural and anthropogenic sources. The natural sources include volcanoes, soils, forests, lakes and open oceans, while the anthropogenic sources are mainly from combustion and waste incineration<sup>[1]</sup>.

In the atmosphere, mercury exists primarily as inorganic forms with two oxidation states: elemental Hg<sup>0</sup> (Gaseous Elemental Mercury, GEM) and divalent Hg(II) (Reactive Gaseous Mercury, RGM, and Particulate Mercury, PHg), Elemental Hg has a long atmospheric lifetime (0.5-2 yrs) and can be transported over a great distance, causing mercury contamination in the remote areas. Divalent Hg, on the other hand, has a short atmospheric lifetime (several hours to days) and deposits back to the earth rapidly through dry and wet depositions due to its strong tendency to be scavenged into droplets or adsorbed onto particulate matter. The atmospheric water also facilitates the wet deposition of mercury, causing contamination of water and soils on earth. The deposited mercury can be reemitted into the atmosphere through a number of chemical and biological pathways leading to the global cycling of mercury<sup>[2]</sup>. Currently, the background total mercury levels are 1–3 ng/m<sup>3</sup>.

Due to the concerns of mercury contamination and its adverse health effects to humans, there is a need to better understand the emission, transport, chemistry and deposition of mercury in the atmosphere. However, the diverse interactions between various mercury species and other atmospheric pollutants are usually complex and nonlinear. Therefore, it requires sophisticated modeling tools to adequately address the fate of atmospheric mercury. In this regard, CMAQ's "one-atmosphere" modeling approach and stateof-the-science components serve as an excellent platform for comprehensive mercury simulation.

CMAQ modeling system has been modified by Bullock and Brehme to simulate the transport, transformation and deposition of atmospheric mercury<sup>[3]</sup>. Although has not been released by CMAS, the latest CMAQ-Hg model is within V 4.3 as a research version. In the model, transformations of Hg are simulated with chemical reactions in the gaseous phase (e.g., Hg reactions with  $O_3$ ,  $CI_2$ ,  $H_2O_2$ , and OH) and in the aqueous phase (e.g., Hg reactions with  $O_3$ , S(IV), HOCI/OCI, HO2, and OH) with a compoundspecific speciation for divalent aqueous Hg using aqueous equilibrium. A linear sorption model is employed to describe the adsorption of dissolved divalent mercury to elemental carbon particles in cloud water. The CMAQ-Hg model simulates the partitioning of RGM between air and cloud water based on the Henry's constant of HgCl<sub>2</sub>. Henry's equilibrium is assumed for GEM. Wet deposition rate is calculated based on precipitation information from MCIP2 and the physicochemical Hg speciation in the cloud chemistry mechanism. Dry deposition rate is calculated based on assumed dry deposition velocity and gaseous concentration for each of the three forms of Hg.

Since the development of the current version of the CMAQ-Hg model, more scientific findings have been reported in the literature. With the new scientific developments in atmospheric mercury, there is a need for further implementations in the model. The objectives of this work are (1) to evaluate the new mercury science components that we recommend to implement in the current version of CMAQ-Hg model and (2) to address the significance of these science implementations. These components are summarized as: (1) improving gaseous-phase mercury chemical mechanism in CMAQ-Hg, (2) implementing sea salt aerosol emission inventory in SMOKE

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processing, (3) implementing reactive halogen activation/reaction mechanism from sea-salt aerosol in CMAH-Hg, (4) implementing new dry deposition velocity schemes for GEM and RGM in MCIP2, and (5) modifying the aqueous divalent mercury sorption algorithms in CMAQ-Hg.

### 2. PROPOSED IMPLEMENTATIONS

2.1 Gaseous-phase Chemical Mechanism

The current version of CMAQ-Hg does not consider the more recent chemical mechanisms for Hg reactions with reactive halogens. These reactions may play an important role for mercury deposition in marine boundary layer, in coastal areas, and near anthropogenic halogen emission sources. Recent modeling study using conditions representing typical marine boundary layer has suggested that the lifetime of elemental mercury in summertime is around 10 days<sup>[4]</sup>. This is much shorter than the generally accepted 0.5 – 2 years atmospheric lifetime of mercury, indicating the potential forcing of mercury deposition by halogen chemistry. These reactions include:

$Hg^{0}_{(g)} + CI_{2(g)} \xrightarrow{\longrightarrow} \text{products}, \ k^{[5]} = 2.6 \times 10^{-18}, \text{ cm}^{3}\text{molec}^{-1}\text{s}^{-1}$
$Hg^{0}_{(g)} + Br_{2(g)} \longrightarrow \text{products},$ $k^{[5]} = 9 \times 10^{-17}, \text{ cm}^{3} \text{molec}^{-1} \text{s}^{-1}$
$Hg^{0}_{(g)} + CI_{(g)} \xrightarrow{\longrightarrow} \text{products}, \ k^{[5]} = 1.0 \times 10^{-11}, \text{ cm}^{3}\text{molec}^{-1}\text{s}^{-1}$
k <sup>[5]</sup> = 1.0×10 <sup>-11</sup> , cm <sup>3</sup> molec <sup>-1</sup> s <sup>-1</sup>
$Hg^{0}_{(g)} + Br_{(g)} \rightarrow \text{products},$ $k^{[5]} = 3.2 \times 10^{-12} \text{ cm}^{3} \text{molec}^{-1} \text{s}^{-1}$
k <sup>[5]</sup> = 3.2×10 <sup>-12</sup> cm <sup>3</sup> molec <sup>-1</sup> s <sup>-1</sup>
$Hg^{0}_{(g)} + BrO_{(g)} \longrightarrow products,$ $k^{[6]} = 1.5 \times 10^{-14} \text{ cm}^{3} \text{molec}^{-1} \text{s}^{-1}$
$k^{[6]} = 1.5 \times 10^{-14} \text{ cm}^3 \text{molec}^{-1} \text{s}^{-1}$

In addition, a recent study on Hg<sup>0</sup>-O<sub>3</sub> kinetics reveals that the rate constant is significantly greater than previously implemented value (7.5×10<sup>-19</sup> vs. 3×10<sup>-20</sup> cm<sup>3</sup>molec<sup>-1</sup>s<sup>-1</sup>)<sup>[7]</sup>. To illustrate the potential impact of these reactions on Hg<sup>0</sup> oxidation, which is the primary driving force for mercury deposition, we computed the oxidation rates contributed by various oxidants under three typical atmospheric scenarios. Table 1 shows the typical concentrations of the oxidants under considerations. With the concentrations in Table 1, the resulted pseudo-first-order oxidation rates of  $Hg^0$  are shown in Figure 1. As seen,  $O_3$  becomes a dominant GEM oxidant in the gaseous phase in all scenarios when the new Hg<sup>0</sup>-O<sub>3</sub> kinetics is implemented. Also, oxidations by CI and Br atoms are important in the Marine Boundary Layer (MBL). By implementing the new gaseous-phase chemical mechanism, mercury simulation in CMAQ should better represent the chemistry of

mercury cycle under various atmospheric conditions. We also recommend testing the impact of changing the  $Hg^0$ - $O_3$  reaction kinetics on mercury deposition in the model.

Table 1. Typical concentrations of Hg <sup>0</sup> oxida	ants
in various scenarios	

Oxidants	Typical Location			Remark		
	Urban	Remote Area	MBL	Remark		
O <sub>3</sub> <sup>(8)</sup> (ppbV)	150	30	30	Daytime		
$O_3^{(8)}$ (molecules cm <sup>-3</sup> )	3.69 x 10 <sup>12</sup>	7.38 x 10 <sup>11</sup>	7.38 x 10 <sup>11</sup>	Daytime		
OH <sup>(9)</sup> (molecules cm <sup>-3</sup> )	5 x 10 <sup>6</sup>	5 x 10⁵	1 x 10 <sup>6</sup>	Daytime		
$H_2O_2^{(10), (11)}$ (molecules cm <sup>-3</sup> )	4.92 x 10 <sup>10</sup>	2.46 x 10 <sup>10</sup>	2.46 x 10 <sup>10</sup>	Daytime		
$Cl_2^{(12), (13)}$ (molecules cm <sup>-3</sup> )	3.69 x 10 <sup>8</sup>	0	1.23 x 10 <sup>9</sup>	Nighttime		
Cl <sup>(5), (12)</sup> (molecules cm <sup>-3</sup> )	1 x 10 <sup>4</sup>	0	5 x 10 <sup>4</sup>	Daytime		
$\operatorname{Br}_{2}^{(14)}$ (molecules cm <sup>-3</sup> )	0	0	2.46 x 10 <sup>7</sup>	Nighttime		
$Br^{(14)}$ (molecules cm <sup>-3</sup> )	0	0	1 x 10⁵	Daytime		
BrO <sup>(14), (15), (16)</sup> (molecules cm <sup>-3</sup> )	0	0	5 x 10 <sup>6</sup>	Daytime		



Figure 1. Comparison of Hg<sup>0</sup> oxidation rates in the scenarios shown in Table 1

#### 2.2 Sea-salt Aerosol Generation

Sea-salt aerosol is the most important source for reactive halogen species in the marine boundary layer. Since reactive halogens play an important role in oxidizing GEM and sea surface cover 71 % of the earth, the inclusion of sea-salt aerosol is critical to simulate the global deposition of mercury. In addition, divalent aqueous mercury can be greatly stabilized in sea salt aerosol due to the strong complex formation with halogen ions, which in turn enhances the deposition of divalent mercury. То realistically represent the concentrations of reactive halogens in coastal areas, marine boundary layer, and remote regions, we recommend including sea salt aerosol

generation in SMOKE processing. The sizedependent generation of sea salt aerosol has been parameterized by Monahan using wind speed at sea surface<sup>[17]</sup>. This parameterization can be built into SMOKE framework to provide sea salt aerosol emission inventory above sea surface.

### 2.3 Reactive Halogen Activation Chemistry

With the emission inventory of sea salt aerosol, the concentration of reactive halogen species can be simulated using a reactive halogen chemical mechanism. activation Reactive halogens can be generated from sea-salt aerosol through acid displacement reactions, oxidation of halides by photochemical oxidants, reactions of chlorine nitrate and bromine nitrates, and even from autocatalytic mechanisms. Several chemical mechanisms of halogen activation from sea-salt aerosol have been developed<sup>[14, 18, 19]</sup> for polluted and remote (clean) marine boundary layers, these mechanisms can be employed to provide the temporally resolved concentrations of important GEM oxidants such as CI, Br, BrO, and HOCI. We recommend incorporating the reactive halogen activation chemistry in CMAQ-Hg for the simulation of mercury-halogen interactions. Although this additional module may increase the computational cost, it is of prime importance to simulate these species for assessing the fate of mercury in the marine boundary layer and in coastal areas.

#### 2.4 Implementation of Dry Deposition Scheme of GEM and RGM

At present, the CMAQ-Hg implements dry deposition velocity for reactive gaseous mercury (RGM) and particulate mercury (PHg). In the treatment of dry deposition, the dry deposition velocity of HNO<sub>3</sub> (ranging from  $0.5 - 8 \text{ cms}^{-1}$  in summer depending on the aerodynamic and boundary layer resistance, from MCIP2 output) is used for the dry deposition of RGM; and the dry deposition velocities of aerosol in Aitken and Accumulation modes are used for the dry deposition velocity of RGM usually ranges from  $0.5 - 2.0 \text{ cm/s}^{[4]}$ , therefore the currently implemented dry deposition velocity may be over-estimating.

Furthermore, the current CMAQ version has no dry deposition scheme for GEM. Dry deposition of GEM may contribute significantly to the total deposition of mercury because of the relatively high concentration of GEM compared to RGM. The dry deposition velocity of GEM estimated from flux and concentration measurements ranges from 0.05 - 0.2 cm/s. Even though the dry deposition velocities of GEM are much lower than those of RGM, the dry deposition flux (the product of concentration multiply by deposition velocity) of GEM and RGM are comparable due to the predominant concentrations of GEM over RGM.

To better simulate total mercury deposition in CMAQ-Hg, we propose to implement the dry deposition velocities for both GEM and RGM in MCIP2:

$$V_d = (r_a + r_b + r_c)^{-1}$$

Where  $r_a$  (aerodynamic resistance) will be estimated from turbulent transport,  $r_b$  (quasilaminar resistance) will be estimated from mercury diffusivities<sup>[20]</sup>, and  $r_c$  (surface resistance) will be estimated from the Henry's constants of GEM and RGM, respectively.

### 2.5 Modification of Aqueous Hg(II) Sorption Algorithm

The current CMAQ-Hg model simulation for the adsorption of aqueous Hg<sup>2+</sup> species to Elemental Carbon Aerosol (ECA) suspended in cloud water is adapted from a sorption study using atmospheric particulate matter (APM) as the sorbent<sup>[3]</sup>. In CMAQ-Hg, the ECA is assumed to be 5 % of the total APM in cloud water. The adsorption is treated as bi-directional nonequilibrium kinetics using a linear aqueous-phase adsorption isotherm for deriving the adsorption ( $k_S$ ) and desorption ( $k_D$ ) rate constants:

$$k_{S} = \frac{[Hg_{S}^{2+}]_{aq}}{t^{*}([Hg_{S}^{2+}]_{aq} + [Hg_{D}^{2+}]_{aq})}$$
$$k_{D} = \frac{[Hg_{D}^{2+}]_{aq}}{t^{*}([Hg_{S}^{2+}]_{aq} + [Hg_{D}^{2+}]_{aq})}$$

Where t is the time constant (3600 s as the time step) for sorption equilibrium,  $[Hg_D^{2+}]_{aq}$  is the dissolved aqueous concentration of Hg(II), and  $[Hg_S^{2+}]_{aq}$  is the adsorbed aqueous concentration of Hg(II). This kinetic treatment may not be optimal since the total Hg(II) available in cloud water is not conserved in the non-equilibrium sorption treatment. Furthermore, the kinetic relationship of sorption and desorption does not include the concentration gradient between the adsorbed and bulk species, which is the fundamental driving force of sorption phenomena.

Since the APM concentration in cloud water is usually small (in the order of magnitude of a few mg/L or lower), and the particles size is usually very small, we recommend directly employing the sorption equilibrium to characterize the quantity of Hg(II) adsorbed onto the ECA particles. This will resolve the mass conservation issues and greatly improve the efficiency of the code in the sorption treatment, i.e.,

 $[Hg_{D}^{2+}]_{total} = (1 + k_{ECA}[ECA])[Hg_{D}^{2+}]_{aq}$ 

Where  $k_{ECA}$  is the sorption equilibrium constant, [*ECA*] is the concentration of ECA in cloud water and  $[Hg_D^{2+}]_{total}$  is the total Hg(II) concentration in bulk and particulate phase in cloud water. We also propose to investigate the impact of the changed sorption algorithm in the wet deposition of mercury in CMAQ-Hg.

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# 4. REFERENCES CITED

<sup>[1]</sup> Mason R. P., Fitzgerald W. F. and Morel F. M. M. (1994): The biogeochemical cycling of elemental mercury: anthropogenic influences. *Geochim. Cosmochim. Acta 58*, pp. 3191-3198.

<sup>[2]</sup> Lin C.-J. and Pehkonen S. O. (1999): Chemistry of atmospheric mercury - A review, *Atmospheric Environment*, *33*, pp. 2067-2079.

<sup>[3]</sup> Bullock O. R. Jr. and Brehme K. A. (2002): Atmospheric mercury simulation using the CMAQ model: formulation description and analysis of wet deposition results, *Atmospheric Environment.* 36, pp. 2135-2146.

<sup>[4]</sup> Hedgecock I.M. and Pirrone N. (2004): Chasing quicksilver: Modeling the atmospheric lifetime of Hg<sup>o</sup><sub>(g)</sub> in the marine boundary layer at various latitudes. *Environ. Sci. Technol. 38*, pp. 68-76.

<sup>[5]</sup> Ariya P.A., Khalizov A., Gidas A. (2002): Reactions of gaseous mercury with atomic and molecular halogens: Kinetics, product studies and atmospheric implications, *J. Phys. Chem. A 106*, pp. 7310-7320.

<sup>[6]</sup> Raofie F. and Ariya P. A. (2004): Product study of the gas-phase BrO-initiated oxidation of Hg0: Evidence for stable Hg<sup>1+</sup> compounds, *Environ. Sci. Technol*, 38,4319-4326. <sup>[7]</sup> Pal B. and Ariya P. A. (2004): Studies of ozone initiated reactions of gaseous mercury: kinetics, product studies, and atmospheric implications, *Phys. Chem. Chem. Phys.*, 6, 572–579.

<sup>[8]</sup> Seinfeld J. H. and Pandis S. N.: *Atmospheric chemistry and physics: From air pollution to climate change*, Wiley-Interscience, 1998, pp. 250, 251, 310.

<sup>[9]</sup> Heard D. E. and Pilling M. J. (2003): Measurement of OH and  $HO_2$  in the troposphere, *Chem. Rev.* 103, 5163-5198.

<sup>[10]</sup> Lee M., Heikes B. G., O'Sullivan D. W. (2000): Hydrogen peroxide and organic hydroperoxide in the troposphere: a review, *Atmospheric Environment*, 34, 3475 – 3494.

<sup>[11]</sup> Reeves C. E. and Penkett S. A. (2003): Measurements of peroxides and what they tell us, *Chemical Reviews*, 103, 5199-5218.

<sup>[12]</sup> Graedel T. E. and Keene W. C. (1995): Tropospheric budget of reactive chlorine, *Global Biogeochemical Cycles*, 9, pp. 47-77.

<sup>[13]</sup> Knipping E. and Dabdub D. (2003): Impact of chlorine emissions from sea-salt aerosol on coastal urban ozone, *Environ. Sci. Technol.* 37, 275 – 284.

<sup>[14]</sup> Von Glasow R. et al., (2002): Modeling halogen chemistry in the marine boundary layer 1. Cloudfree MBL, *Journal of Geophysical Research*, vol. 107, no. D17, 4341.

<sup>[15]</sup> Stutz J. and Ackermann R. (2002): Atmospheric reactive chlorine and bromine at the Great Salt Lake, Utah, *Geophysical Research Letters*, 29, 1380.

<sup>[16]</sup> Wayne et al., (1995): Halogen Oxides: radicals, sources and reservoirs in the Laboratory and in the Atmosphere, *Atmospheric Environment*, 29, pp. 2677-2881.
<sup>[17]</sup> E.C. Monahan and G. MacNiociall: *Oceanic*

[<sup>17]</sup> E.C. Monahan and G. MacNiociall: Oceanic whitecaps and their role in air-sea exchange processes, 1986, D. Reidel Publishing, 167-174.
 [<sup>18]</sup> Vogt R., Crutzen P. J. and Sander R. (1996):

<sup>[18]</sup> Vogt R., Crutzen P. J. and Sander R. (1996): A mechanism for halogen release from sea-salt aerosol in the remote marine boundary layer, *Nature*, 382, 327-330.

<sup>[19]</sup> Knipping E. M. and Dabdub D. (2002).Modeling Cl<sub>2</sub> formation from aqueous NaCl particles:

Evidence for interfacial reactions and importance of Cl<sub>2</sub> decomposition in alkaline solution, *Journal of Geophysical Research*, 107, No. D18, 4360,

<sup>[20]</sup> Massman W. J. (1999): Molecular diffusivities of Hg vapor in air,  $O_2$  and  $N_2$  near STP and the kinematic viscosity and thermal diffusivity of air near STP, *Atmospheric Environment*, 33 (1999) 453-457