

## THE SENSITIVITY OF AEROSOL SULFATE TO CHANGES IN NITROGEN OXIDES AND VOLATILE ORGANIC COMPOUNDS

Ariel F. Stein \*

Department of Meteorology, The Pennsylvania State University

e-mail: [ais5@psu.edu](mailto:ais5@psu.edu)

Voice (301) 713-0295 x 119

Fax (301) 713-0119

### 1.0 INTRODUCTION

Traditionally, strategies to control sulfate formation have been focused on the response of sulfate ( $\text{SO}_4^{2-}$ ) to reductions in sulfur dioxide ( $\text{SO}_2$ ) emissions [e.g. Shin and Carmichael, 1992]. Much less attention has been devoted to assessing the effects on sulfate formation arising from changes in emissions of volatile organic compounds (VOC) and nitrogen oxides ( $\text{NO}_x$ ), the chemical families largely responsible for the generation of the tropospheric ozone ( $\text{O}_3$ ) and the oxidants needed to transform  $\text{SO}_2$  to  $\text{SO}_4^{2-}$ . The response of ambient sulfate to reductions in  $\text{NO}_x$  and VOC emissions depends in part on the resulting changes in oxidant levels and the competition that naturally exists between in-cloud (*via* hydrogen peroxide) and clear-air (*via* hydroxyl radical)  $\text{SO}_2$  oxidation. Coupling of the different chemical families leads to complex and potentially non-linear behavior with non-intuitive consequences (Stein and Lamb, 2000).

### 2.0 $\text{SO}_4^{2-}$ SENSITIVITY TO $\text{NO}_x$ AND VOC

The idea of "potential" sulfate can be utilized to study the combined response of aqueous- and gas-phase sulfate formation to changes in  $\text{NO}_x$  and VOC emissions using one conceptual species only. Potential sulfate, defined as the sum of ambient sulfate (i.e. total sulfate) and hydrogen peroxide ( $\text{H}_2\text{O}_2$ ) [Stockwell, 1994], can be interpreted as the maximum concentration of  $\text{SO}_4^{2-}$  that could be produced in the gas and aqueous phases. The usefulness of potential sulfate as a surrogate for ambient sulfate depends on the prevailing meteorological conditions and emission distributions. Nevertheless, this concept provides an integrated way to deal with this complex chemical system.

We can identify two photochemical regimes, each of which exhibits a different response of potential  $\text{SO}_4^{2-}$  production to changes in  $\text{NO}_x$  and VOC emissions, depending on the main sink of odd-hydrogen ( $\text{HO}_x$ ), defined as the sum of hydroxyl (OH), hydroperoxyl ( $\text{HO}_2$ ), and organic peroxy ( $\text{RO}_2$ ) radicals [Sillman, 1995]. When the formation of nitric acid ( $\text{HNO}_3$ )

dominates the loss of odd-hydrogen, a decrease in  $\text{NO}_x$  frees up OH that can react with  $\text{SO}_2$ , CO, and VOC, which thus increases the abundance of  $\text{HO}_2$  and in turn favors the formation of  $\text{H}_2\text{O}_2$ . Consequently, the interaction of these reactions tends to increase the formation of  $\text{SO}_4^{2-}$  via both the clear-air and aqueous-phase pathways. A decrease in VOC levels would reduce OH by making it available to form more  $\text{HNO}_3$ . As a consequence, the  $\text{H}_2\text{O}_2$  production would also be reduced, thus decreasing the formation of  $\text{SO}_4^{2-}$ . Under this photochemical regime, called "VOC-sensitive", potential  $\text{SO}_4^{2-}$  decreases with decreasing VOC and increases with decreasing  $\text{NO}_x$ .

By contrast, when the formation of  $\text{HNO}_3$  can be regarded as a small odd-hydrogen sink, a decrease in  $\text{NO}_x$  concentrations would slow down the conversion of  $\text{HO}_2$  to OH, thereby decreasing the overall  $\text{SO}_4^{2-}$  formation. When the VOC levels are lowered, a small additional amount of OH would be available to react with  $\text{SO}_2$ , slightly increasing the gas-phase production of  $\text{SO}_4^{2-}$ . Lowering the VOC concentrations would not significantly affect the formation of  $\text{H}_2\text{O}_2$ , however, due to the fact that the system is saturated with hydrocarbons. Under this condition, changes in VOC levels would have little effect on  $\text{HO}_2$  and hence on  $\text{H}_2\text{O}_2$  formation [Sillman *et al.*, 1990]. Therefore, the aqueous-phase  $\text{SO}_2$  oxidation would remain unaltered. This scenario constitutes the " $\text{NO}_x$ -sensitive" photochemical regime, in which aerosol  $\text{SO}_4^{2-}$  concentrations would decrease with decreasing  $\text{NO}_x$  and would be largely independent of VOC. The highly nonlinear and coupled nature of the sulfate production process is responsible for this somewhat counter-intuitive behavior.

Milford *et al.* [1994] and Sillman [1995] have developed a methodology to assess the sensitivity of a secondary pollutant (ozone) to changes in  $\text{NO}_x$  and VOC. They found that model predictions for  $\text{O}_3$ - $\text{NO}_x$  sensitivity were associated with the concentrations of a number of key chemical species that differed from those linked to  $\text{O}_3$ -VOC sensitivity. Milford *et al.* [1994] showed that ozone simulated under an  $\text{NO}_x$ -sensitive regime was linked to afternoon values of total reactive nitrogen ( $\text{NO}_y = \text{NO}_x + \text{HNO}_3 + \text{Peroxyacetylnitrates} + \text{alkyl nitrates}$ ) below a certain threshold concentration, while VOC-sensitive ozone was connected to higher  $\text{NO}_y$  levels. Sillman [1995] extended that work to include species ratios, such as  $\text{O}_3/\text{NO}_y$  and  $\text{H}_2\text{O}_2/\text{HNO}_3$ . Following these ideas, we propose the use of a combination of afternoon concentrations of  $\text{HNO}_3$ ,  $\text{H}_2\text{O}_2$ ,

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\* Corresponding author address: Ariel F. Stein, ARL – NOAA, 1315 East West Hwy, Silver Spring, MD 20910.

and ambient  $\text{SO}_4^{2-}$  as “indicator species” of the ambient  $\text{SO}_4^{2-}$ -VOC- $\text{NO}_x$  sensitivity.

The link between the indicator value and the  $\text{NO}_x$ -VOC chemistry can be understood in terms of the dominant sinks and sources for odd-hydrogen. The  $\text{NO}_x$ -sensitive regime favors the formation of potential sulfate over the production of  $\text{HNO}_3$ . Under this condition the production of  $\text{H}_2\text{O}_2$  constitutes the main loss of  $\text{HO}_x$ . Also, under  $\text{NO}_x$ -sensitive conditions the gas-phase  $\text{SO}_2$  oxidation is favored over the formation of  $\text{HNO}_3$ . On the other hand, the VOC-sensitive regime is characterized by a high production rate of  $\text{HNO}_3$  that overwhelms the formation of potential sulfate. Therefore, high values of the indicator ratio,  $\frac{[\text{H}_2\text{O}_2]+[\text{SO}_4^{2-}]}{[\text{HNO}_3]+[\text{NO}_3]}$ , are associated with  $\text{NO}_x$ -sensitive conditions while low values for the indicator can be identified with a VOC sensitive regime.

### 3.0 SIMULATION

In order to investigate the response of “potential” sulfate to changes in the  $\text{NO}_x$  and VOC source strengths we make use of U.S. Environmental Protection Agency (EPA)’s MODELS-3 [Dennis *et al.*, 1996] Personal Computer (PC) version. This system constitutes a state-of-the-art tool for regional-scale simulations of photochemical smog, visibility and fine particulates. MODELS-3 is a versatile software system that includes a Community Multiscale Air Quality (CMAQ) model [EPA, 1999], a three-dimensional Eulerian chemical transport model that accounts for horizontal and vertical advection, eddy diffusion, gas-phase chemical transformations, emissions, cloud mixing, aqueous-phase chemical reactions, and aerosol processes. The meteorological data used by this simulation are obtained as output fields from the Mesoscale Model Version 5 (MM5) [Grell *et al.*, 1994, Seaman and Michelson, 2000]. Emissions are calculated by the MODELS-3 Emissions Processing and Projection System (MEPPS) [EPA, 1999]. A detailed description of the equations and algorithms of each component of MODELS-3 are given by EPA [1999].

The simulation results presented here are based on an air pollution event that took place over the eastern United States during July 12-14, 1995. This episode featured concentrations exceeding the 1-h National Ambient Air Quality Standards (NAAQS) for ozone with a maximum of 175 parts per billion (ppb) measured at a Connecticut site on July 14 [Sistla *et al.*, 2001]. In order to allow the model to build up the chemical species concentrations, thus minimizing the influence of initial conditions, the first two simulation days were treated as an initialization period. Therefore, only the results for July 14 are discussed here. The model domain covered an area of 2268 x 2808  $\text{km}^2$  from the Great Lakes and northern New England to the Gulf of Mexico with a 36-km resolution. The simulation utilized 15 non-hydrostatic sigma-pressure vertical layers. Initial concentrations and boundary conditions were set to near-rural conditions [Chang *et al.*, 1990]. The RADM2 chemical mechanism [Stockwell *et al.*, 1990] was used to simulate the gas-phase chemical reactions. The aerosol processes were

calculated based on the Regional Particulate Model (RPM) [Binkowski and Shankar, 1995]. Area and point emissions were estimated using MEPPS, which is based on the 1995 National Emissions Trends [EPA, 1999]. Emissions data from motor vehicles and vegetation were simulated using the Mobile5a [EPA, 1996] and BEIS-2 [Pierce *et al.*, 1998] models, respectively.

The meteorological data used for the present case study have been described in detail in Seaman and Michelson [2000], so only a brief summary is given here. This event featured light winds, restricted vertical mixing and high temperatures, conditions typically associated with a Bermuda high. Mesoscale structures associated with an Appalachian lee trough played a key role in determining the geographical distribution of photochemical smog. Low mixing depths and south-south-westerly flow favored the accumulation of photochemical precursors to the east of the trough [Seaman and Michelson, 2000], further explaining the observed ozone concentrations in excess of 120 ppb along the northeastern U.S. urban corridor.

The data used to verify the model predictions are a subset of measurements made during the North American Research Strategy for Tropospheric Ozone-Northeast (NARSTO-NE) 1995 intensive field campaign [Mueller, 1998]. A total of 327 measurement sites distributed throughout the Northeastern United States were selected to make the model comparison. The modeled and measured ozone concentrations were paired spatially by performing a bi-linear horizontal interpolation of the simulated values to the corresponding monitor locations.

### 4.0 RESULTS AND DISCUSSION

Good agreement is found between modeled and observed maximum 1-h ozone concentrations for the Northeastern U.S. The simulation captures the broad features of the geographical distribution of this episode, but not the detailed spatial variability. In particular, the highest ozone peaks are under-predicted by the model. The simulation shows a raw bias of -9.7 ppb, a normalized bias of -7.3%, a raw gross error of 18.3 ppb, and a normalized gross error of 18.5% for the spatially paired maximum 1-h ozone mixing ratios. These simulation results are within the performance “goals” suggested by Tesche *et al.* [1990] and by Hanna *et al.* [1996].

In order to build further confidence in the ability of the model to predict sensitivities of photochemical species to changes in VOC and  $\text{NO}_x$ , we performed three separate series of runs, each with 35% reductions in emission rates separately for VOC and  $\text{NO}_x$ . The first simulation (BASECASE) corresponds to the base case scenario with standard emission rates. In the second run (DBLSO2) the  $\text{SO}_2$  emissions rates were doubled from those of the base case. In the third scenario (DBLVOC) the anthropogenic VOC emission rates were doubled with respect to the base case. Subsequently, we calculated the values for two concentration ratios, namely  $\text{O}_3/\text{NO}_2$  ( $\text{NO}_2 = \text{NO}_y\text{-NO}_x$ ) and  $\text{H}_2\text{O}_2/\text{HNO}_3$ , that

are associated with ozone- NO<sub>x</sub>- and VOC-sensitive locations at 16:00 EDT (20:00 UTC). VOC-sensitive locations are defined as those where the O<sub>3</sub> concentration in the simulation with reduced VOC is lower than the levels of O<sub>3</sub> in both the base case scenario and in the simulation with an equivalent reduction in NO<sub>x</sub> by at least 5 ppb [Sillman *et al.*, 1997]. NO<sub>x</sub>-sensitive locations are classified analogously. The median values along with the 5<sup>th</sup> and 95<sup>th</sup> percentiles for the indicator ratios corresponding to NO<sub>x</sub>-sensitive and VOC-sensitive locations are not substantially different from those calculated by Sillman [1995; 1997]. NO<sub>x</sub>-sensitive locations are associated with O<sub>3</sub>/NO<sub>z</sub> > 12 and H<sub>2</sub>O<sub>2</sub>/HNO<sub>3</sub> > 0.5, while VOC-sensitive stations are linked to O<sub>3</sub>/NO<sub>z</sub> < 9 and H<sub>2</sub>O<sub>2</sub>/HNO<sub>3</sub> < 0.3.

Following the methodology developed by Sillman [1995], we used the  $\frac{[H_2O_2]+[SO_4^{2-}]}{[HNO_3]+[NO_3^-]}$  ratio as an indicator for the sensitivity of potential sulfate to changes in NO<sub>x</sub> and VOC emissions. Figure 1 depicts the potential SO<sub>4</sub><sup>2-</sup>-VOC-NO<sub>x</sub> sensitivity for the BASECASE simulation at 16:00 EDT (20:00 UTC). This figure shows the percentage normalized reduction in  $\{[H_2O_2]+[SO_4^{2-}]\}$  concentration as a consequence of either a NO<sub>x</sub> or a VOC emission reduction for each grid within the model domain. The change in potential sulfate is plotted as a function of the concurrent indicator ratio,  $\frac{[H_2O_2]+[SO_4^{2-}]}{[HNO_3]+[NO_3^-]}$ . As can be inferred, there is well-defined contrast between NO<sub>x</sub> and VOC sensitive locations. High  $\frac{[H_2O_2]+[SO_4^{2-}]}{[HNO_3]+[NO_3^-]}$  ratios are associated with reductions in  $\{[H_2O_2]+[SO_4^{2-}]\}$  concentrations as NO<sub>x</sub> emissions are reduced, while no sensitivity is observed for changes in VOC (NO<sub>x</sub> sensitive regime). On the other hand, low indicator ratios are linked to a decrease in "potential" sulfate as the VOC emissions are reduced and an *increase* in the  $\{[H_2O_2]+[SO_4^{2-}]\}$  concentrations when the NO<sub>x</sub> emissions decrease (VOC sensitive regime).

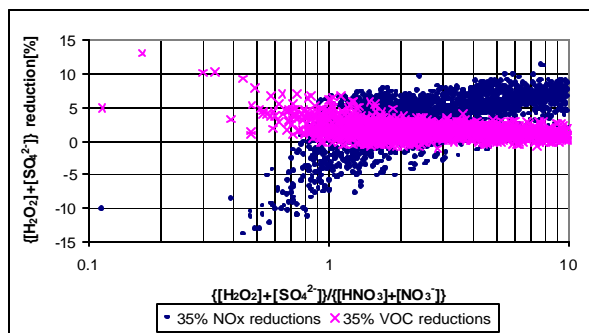


Fig 1 Normalized percentage response of potential sulfate concentrations to changes in NO<sub>x</sub> and VOC versus  $\frac{[H_2O_2]+[SO_4^{2-}]}{[HNO_3]+[NO_3^-]}$  ratios for 20:00 UTC, July 14, 1995.

The correspondence between the SO<sub>4</sub><sup>2-</sup>-VOC-NO<sub>x</sub> sensitivity and the indicator ratio was evaluated quantitatively by calculating the 95<sup>th</sup> and 5<sup>th</sup> percentiles for the  $\frac{[H_2O_2]+[SO_4^{2-}]}{[HNO_3]+[NO_3^-]}$  ratios

associated with VOC- and NO<sub>x</sub>-sensitive locations respectively (Table 1). We define NO<sub>x</sub>-sensitive locations as those where the normalized  $\frac{[H_2O_2]+[SO_4^{2-}]}{[HNO_3]+[NO_3^-]}$  calculated in the simulation with a 35% reduction in NO<sub>x</sub> is lower than the normalized potential sulfate concentrations modeled with a 35% reduction in VOC by at least 5%. VOC-sensitive stations are defined in an analogous way. Table 1 also shows that the median values of the  $\frac{[H_2O_2]+[SO_4^{2-}]}{[HNO_3]+[NO_3^-]}$  ratios linked to NO<sub>x</sub>-sensitive stations are at least four times higher than those associated with VOC-sensitive locations. The 95<sup>th</sup> percentile of the collection of indicator values corresponding to VOC-sensitive stations, which identifies the highest values for this distribution, along with the 5<sup>th</sup> percentile of indicator values associated with NO<sub>x</sub>-sensitive locations, reflecting the lowest values of the indicator distribution, identify the threshold interval of the  $\frac{[H_2O_2]+[SO_4^{2-}]}{[HNO_3]+[NO_3^-]}$  ratio for the transition from VOC to NO<sub>x</sub> sensitivity. The choice of these percentile cutoffs are based on Sillman *et al.* (1997). We find similar transition points for the three different model scenarios that range between 1.4 to 2.2.

Run	VOC sensitive locations		NO <sub>x</sub> sensitive locations	
	50 <sup>th</sup>	95 <sup>th</sup>	5 <sup>th</sup>	50 <sup>th</sup>
	percentile	percentile	percentile	percentile
BASECASE	0.96	1.82	1.54	7.30
DBLSO2	0.89	1.37	1.73	4.76
DBLVOG	1.23	2.25	1.60	7.46

Table 1 Distribution of  $\frac{[H_2O_2]+[SO_4^{2-}]}{[HNO_3]+[NO_3^-]}$  ratios for NO<sub>x</sub>- and VOC- sensitive regimes.

## 5.0 CONCLUSIONS

MODELS-3 constitutes a suitable and powerful tool to assess the sensitivity of particulate sulfate to changes in primary pollutant source strength. In particular, the simulation results show that the response of potential sulfate levels to changes in VOC and NO<sub>x</sub> emissions is strongly correlated with afternoon values of the  $\frac{[H_2O_2]+[SO_4^{2-}]}{[HNO_3]+[NO_3^-]}$  indicator ratio. The relatively narrow range of the transition values unequivocally identifies either a VOC- or NO<sub>x</sub>-sensitive regime despite the three different scenarios used in this modeling exercise. Under VOC-sensitive conditions (when the indicator ratio is less than approximately 1.4) potential sulfate is reduced when VOC emissions are reduced, but it is *increased* when NO<sub>x</sub> emissions are reduced. On the other hand, under NO<sub>x</sub>-sensitive conditions (values higher than 2.2 for the non-dimensional indicator) potential sulfate is reduced when NO<sub>x</sub> emissions are reduced, while it is only weakly sensitive to changes in VOC.

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